

Lecture notes

"Sun-planet connections"

given at MPS

23 - 25 Feb. 2009

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Magnetic field in electrically conductive medium (fluid, gas)

Electric field

$$\vec{E} = -\underbrace{\vec{V} \times \vec{B}}_{\text{convective}} + \frac{1}{\sigma} \underbrace{\vec{j}}_{\text{Ohmic}} \quad \dots (1)$$

$$= -\vec{V} \times \vec{B} + \frac{1}{\mu_0 \sigma} \nabla \times \vec{B} \quad \dots (2)$$

(Ampère's law)

Induction eq.

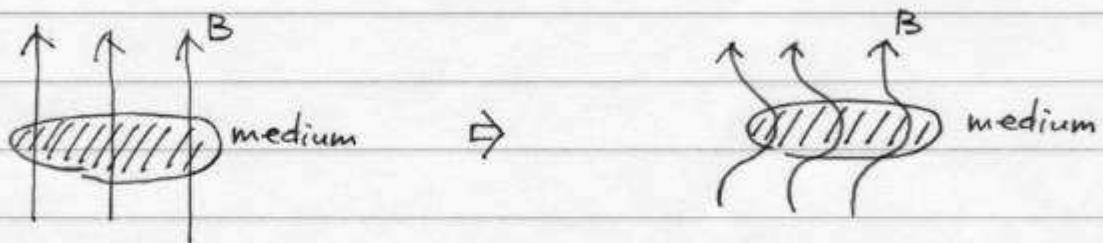
$$\partial_t \vec{B} = -\nabla \times \vec{E} \quad \dots (3)$$

Combine (2) and (3) and $\nabla \cdot \vec{B} = 0$

↗

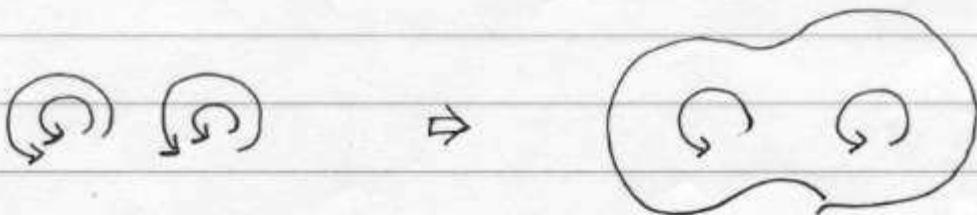
$$\partial_t \vec{B} = \underbrace{\nabla \times (\vec{V} \times \vec{B})}_{\text{Evolution of mag. field}} + \underbrace{\frac{1}{\mu_0 \sigma} \nabla^2 \vec{B}}_{\text{(A) frozen-in (B) diffusion}} \quad \dots (4)$$

(A) Frozen-in magnetic field



Magnetic field moves with flow/medium.

(B) Diffusion



Magnetic field becomes weaker.

Space plasma ... collisionless, conductive medium.
Conductivity σ very large.

↗ Frozen-in magnetic field

Consequences

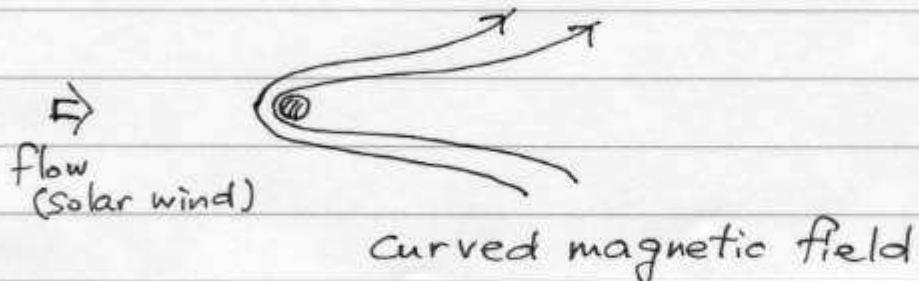
(1) Magnetic field moves with flow
(e.g. solar wind)

(2) Different media cannot be mixed.

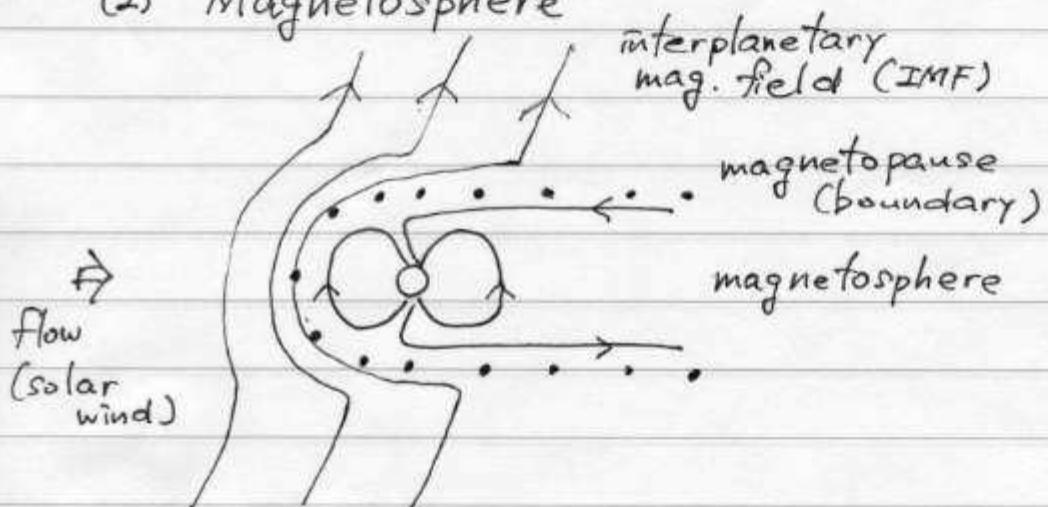
↗ Formation of boundaries
(e.g. magnetopause)

Applications

(a) Comet tail (Alfvén's idea, 1957)



(2) Magnetosphere



How large is magnetosphere?

Pressure balance

$$\frac{1}{2} \rho v^2 = \frac{B^2}{2\mu_0} \quad \dots (5)$$

dynamic pressure

(kinetic energy density)

magnetic pressure

(mag. energy density)

Solar wind

planet

Dipole magnetic field (for planet)

$$B \propto \frac{1}{r^3}$$

$$\therefore \frac{B}{B_{\text{surf}}} = \left(\frac{R_{\text{surf}}}{R_{\text{mp}}} \right)^3 \quad \dots (6)$$

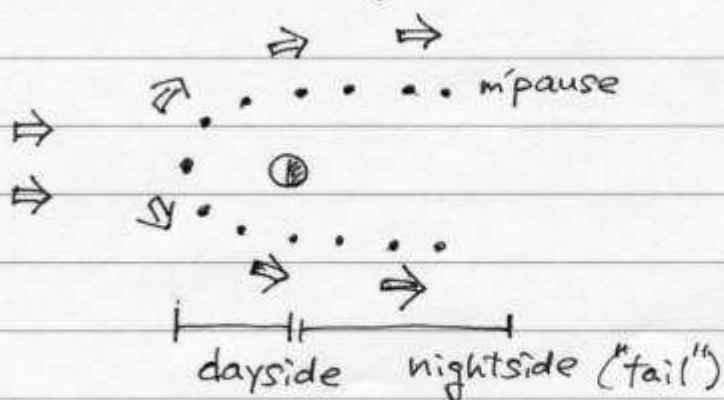
Combine (5) and (6)

$$\therefore \frac{R_{\text{mp}}}{R_{\text{surf}}} = \left(\frac{B_{\text{surf}}^2}{\mu_0 \rho V^2} \right)^{1/6} \quad \dots (7)$$

Classification of planets

	$R_{\text{mp}} > R_{\text{surf}}$	$R_{\text{mp}} < R_{\text{surf}}$
Solid Surface	Mercury	Earth Moon
gas surface (atmosphere)	Earth Jupiter, Saturn Uranus, Neptune	Venus, Mars

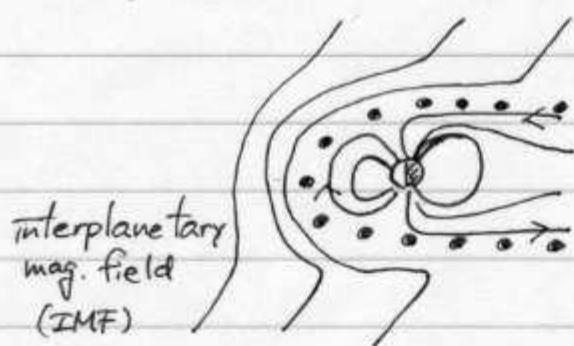
- Flow around magnetosphere



Surface mag. field 30 000 nT

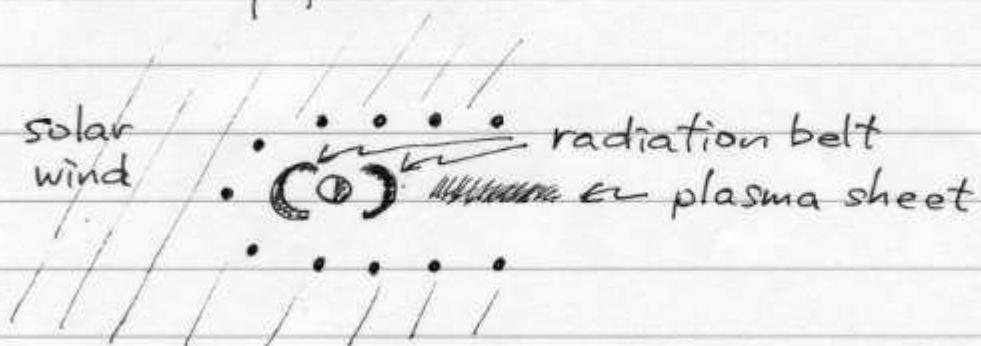
Magnetopause distance ... 10 ~ 11 R_E (dayside)

- Magnetic field

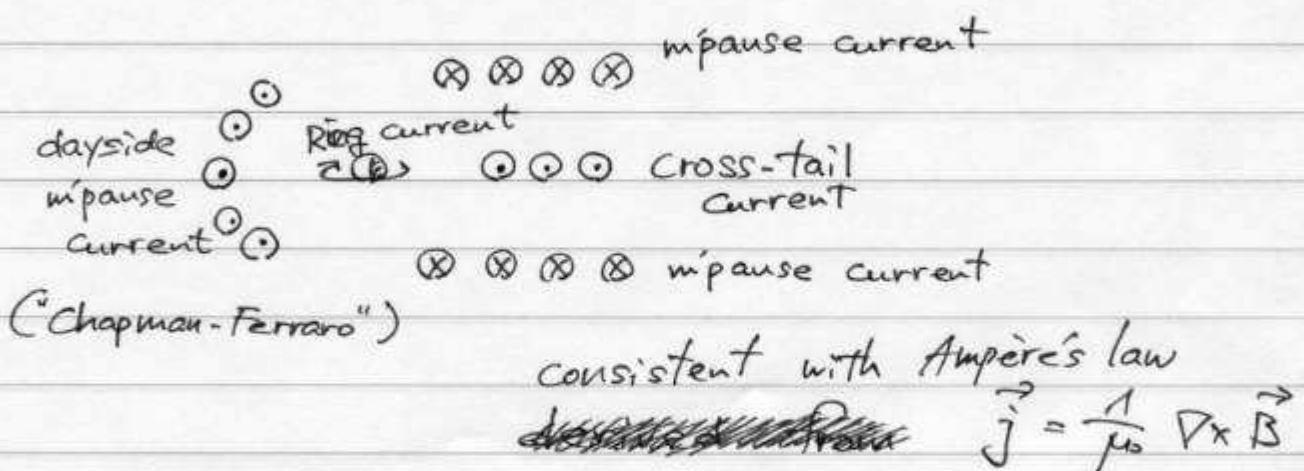


solar wind ... curved mag. field
dayside magnetosphere
... dipolar field
nightside ... elongated / stretched
field

- Plasma populations

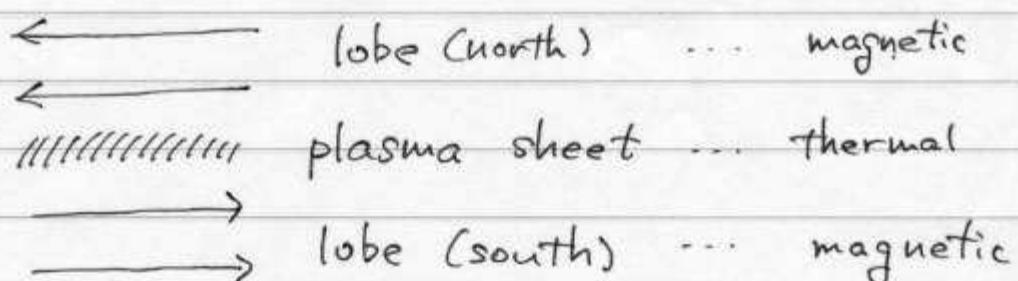


- Electric current



- Pressure balance in tail

pressure

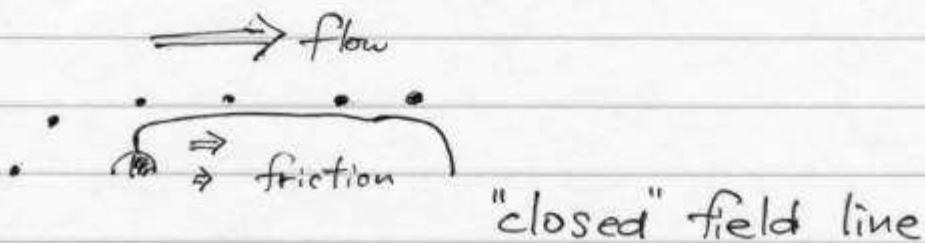


$$\frac{P_{th}}{\text{plasma sheet}} = \frac{P_{mag}}{\text{lobe}}$$

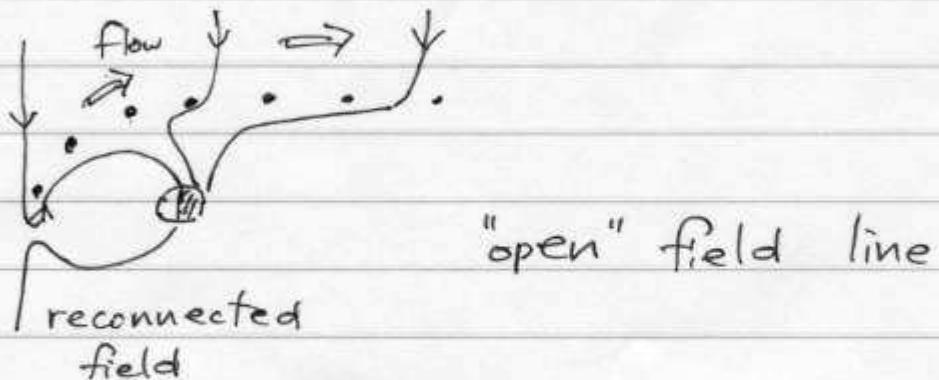
- Why is tail-field stretched?

possibility 1 ... viscosity / friction

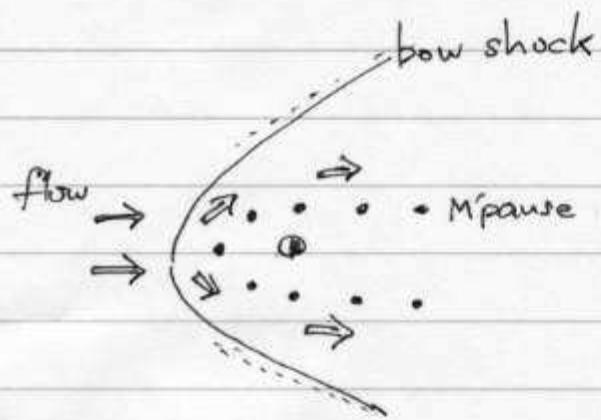
(Axford and Hines, 1961)



possibility 2 ... reconnection (Dungey 1961)



- bow shock



- standing shock wave
- misphere as obstacle
- flow speed changes from super-magnetosonic to sub-magnetosonic

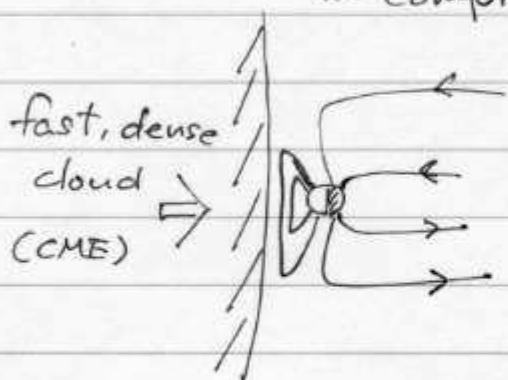
cf. magnetosonic speed

$$V \sim \sqrt{V_A^2 + C_s^2}$$

- plasma is
 - ⁽¹⁾ compressed,
 - ⁽²⁾ heated
- magnetic field also compressed

Geomagnetic storm

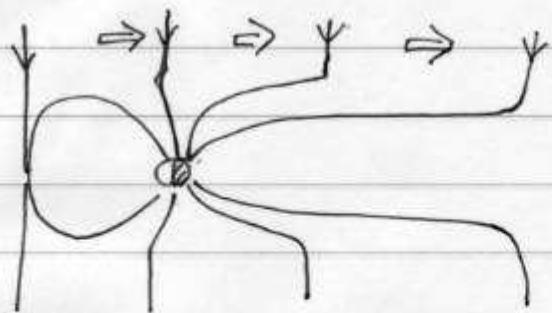
... compression of magnetosphere
(Chapman, 1930)



1. dayside compression ~ minutes
2. nightside compression ~ hours
3. recovery phase (expansion) ~ days

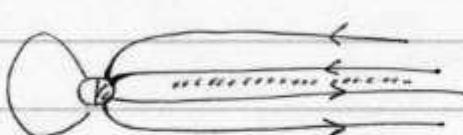
Substorm (reconnection model)

1. growth phase ($T \sim 40$ min.)



- magnetic field transport from dayside to tail (energy charge in tail)
- reconnection on dayside (southward IMF)

2. expansion phase ($T \sim 10 \text{ min.}$)



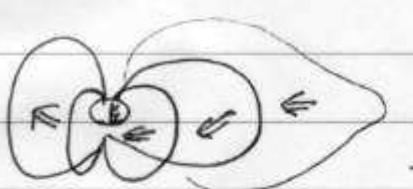
thin current sheet
anti-parallel mag. field

↓ tail reconnection (energy discharge)



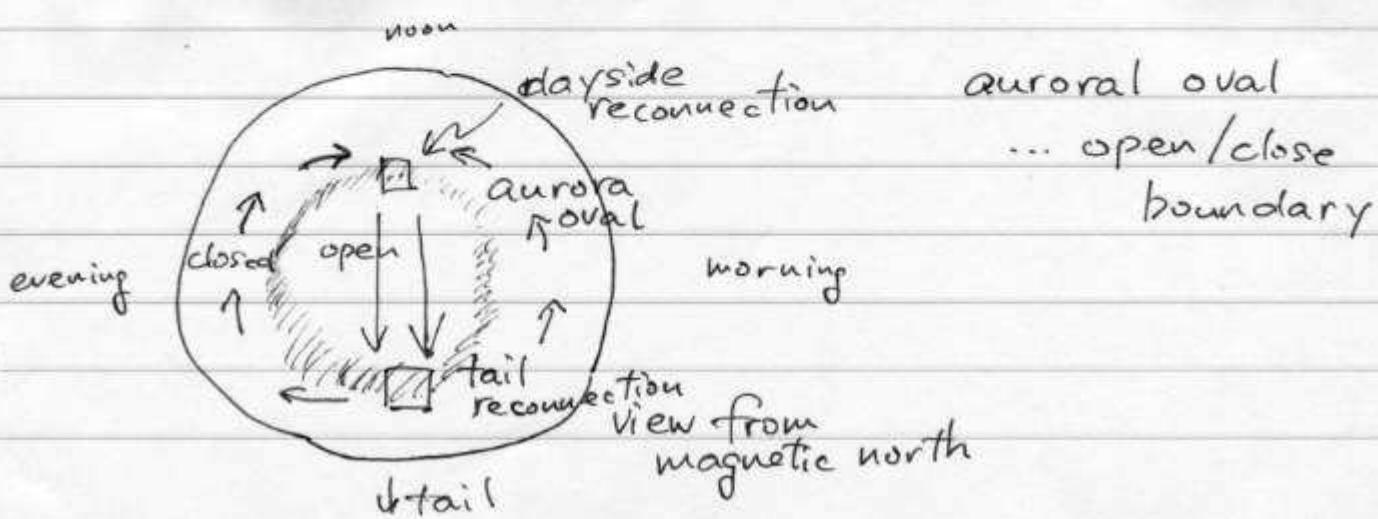
- "dipolarization"
- plasma jet ("plasmoid")
- field-aligned current
into/out of ionosphere
→ bright aurora

3. recovery phase ($T \sim \text{hours}$)



- magnetic field transport
back to dayside

ionospheric convection pattern (Southward IMF)



Chap. 4 Mercury

4.1

Magnetized, non-atmospheric body.

Surface magnetic field $\sim 200 \text{ nT}$

(Mariner 10, Messenger)

cf. 30000 nT at Earth surface

Magnetopause distance

$$\frac{1}{2} \rho v^2 = \frac{B^2}{2\mu_0} \quad (\text{pressure balance})$$

$$B \propto \frac{1}{r^3} \quad (\text{dipole field})$$

$$B_0 = 200 \text{ (nT)} \quad (\text{surface})$$

$$\rho = m_p \times 20 \text{ (cm}^{-3}\text{)} \quad (\text{solar wind mass density})$$

$$v = 300 \text{ (km/s)} \quad (\text{solar wind speed})$$

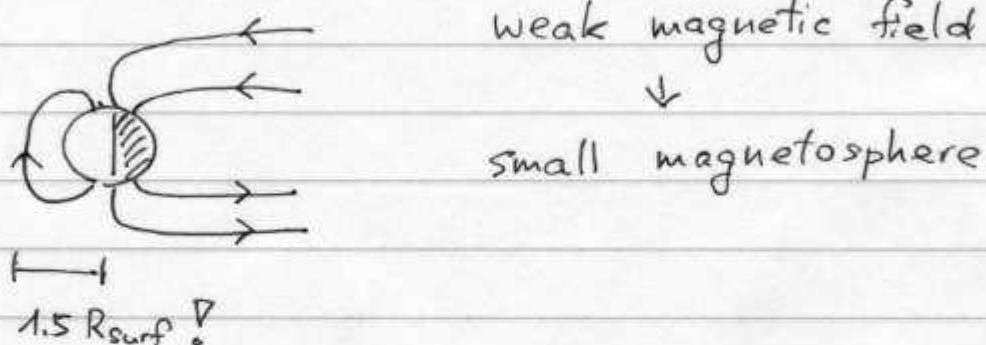
$$\Downarrow R_{mp} \approx 1.5 R_{surf} \quad !$$

(from planet center)

surface - magnetopause distance

$$0.5 R_{surf} = 1200 \text{ km}$$

Magnetosphere



- No radiation belt (or ring current)
... planet ~~occupies~~ occupies a lot of volume
- Sodium (Na) sputtering
→ sodium-rich plasma
- "Hybrid" scale ... magnetosphere size of the order of ion gyro-radius
... → currents are carried by electrons? (mpause, plasma sheet)
- Fast reaction, rapid reconfiguration

length dayside-tail $\ell \sim 10 R_{\text{surf}}$ (24000 km)
solar wind speed $v \sim 300 \text{ km/s}$
Δ
reaction time $\tau = \ell/v \sim 80 \text{ s}$
(cf. 40 min. at Earth)
- reconnection? substorm?

Dynamo problem

Theoretical estimate of surface magnetic field

$$\text{Elsässer number} = \frac{|\text{Lorentz force}|}{|\text{Coriolis force}|}$$

$$= \frac{|\sigma(\vec{J} \times \vec{B}) \times \vec{B}|}{|2\mu_0 \vec{\Omega} \times \vec{v}|}$$

$$\sim \frac{\sigma B^2}{2\mu_0 \Omega}$$

$$\text{Assume } \Lambda \sim 1, \quad B \propto \sqrt{\Omega}.$$

Scale Earth magnetic field into Mercury,

$$B(\text{earth}) = 30000 \text{ nT}$$

$$\Omega(\text{earth}) = 2\pi/1 \text{ (rad/day)}$$

$$\Omega(\text{merc}) = 2\pi/58 \text{ (rad/day)}$$

$$\therefore B(\text{merc}) = 4000 \text{ nT.}$$

Mercury field is weaker than
Earth-scaled field.

Lost dynamo

Venus and Mars have atmosphere
but no global magnetic field.
(dipole)

Estimate of dipole moment (upper limit!)

Venus $10^{-5} \times$ (Earth dipole)
 $\approx 0.1 \text{ nT}$ at surface PVO

Mars $10^{-5} \times$ (Earth dipole)
 $\approx \frac{1}{0.5} \text{ nT}$ at surface MGS

Mars have magnetized crust
e.g. 200 nT at 400 km altitude

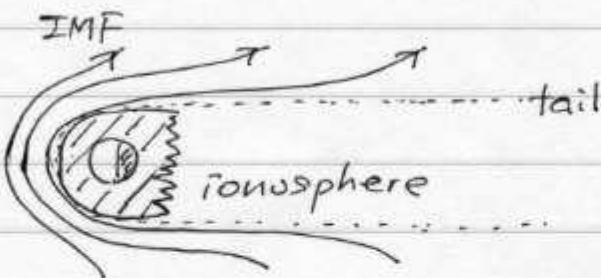
→ Dynamo operating in early time?

Why is dynamo missing?

Venus ... slow rotation (293 days)?

Mars ... core in solid state?
(Liquid core became solid.)
Small planet size

Magnetosphere



Pressure balance

(1) magnetic boundary

$$\frac{1}{2} \rho v^2 = \frac{B^2}{2\mu_0}$$

solar wind
dynamic
pressure

solar wind
magnetic pressure

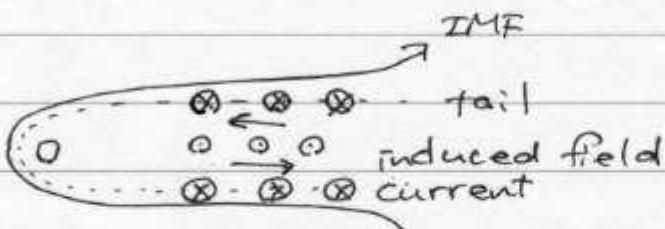
field compression
on dayside

(2) ionopause

$$\frac{1}{2} \rho v^2 = nkT$$

ionosphere
gas pressure

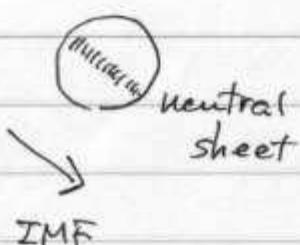
Induced magnetic field



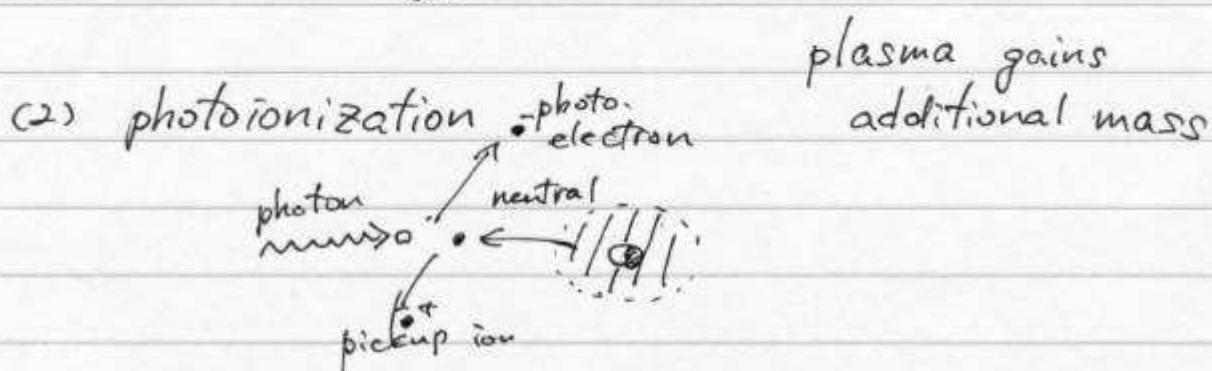
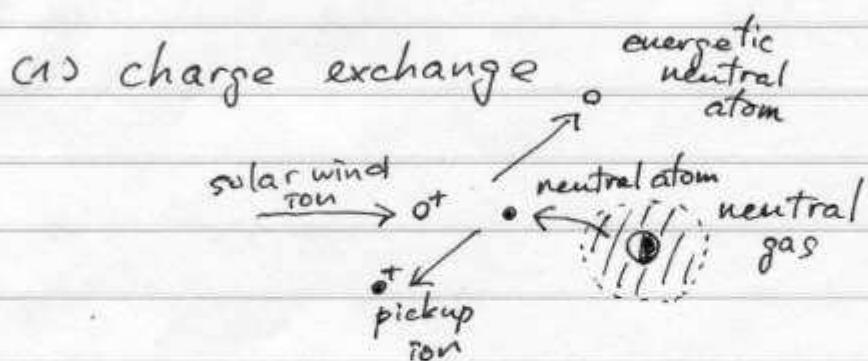
⇒ "induced magnetosphere"

sensitive to IMF direction

tail cross section



Pickup process



Giant magnetosphere

surface magnetic field

400 000 nT (Jupiter)

20 000 nT (Saturn)

metallic hydrogen core?

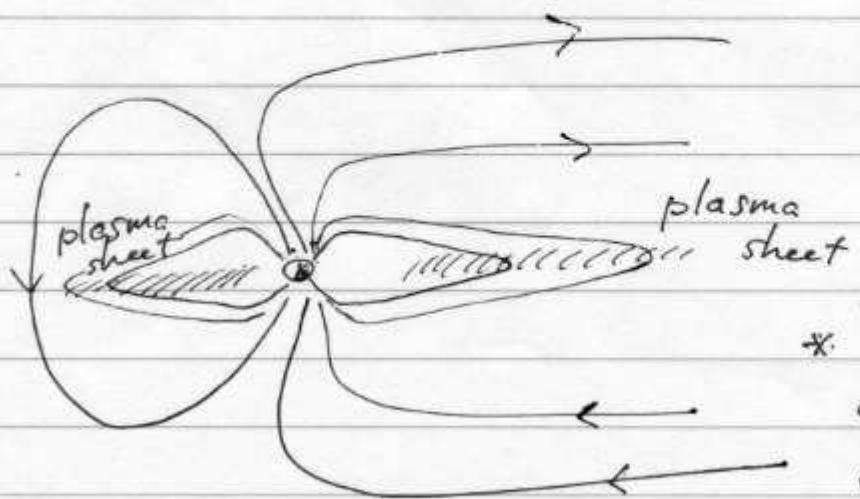
magnetopause distance

50 - 100 R_{surf} (Jupiter) ca. 5 R_{sun}!20 R_{surf} (Saturn) ca. 2 R_{sun}

rotation-dominant

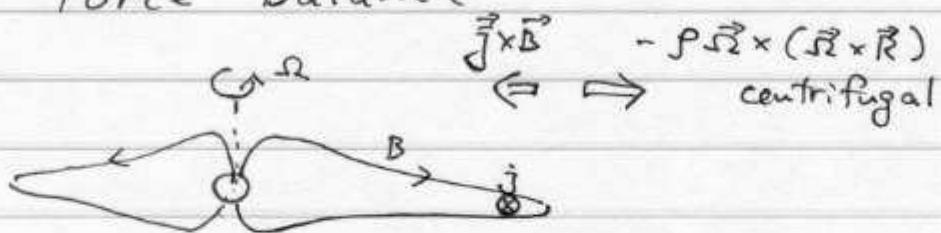
10 hours (Jup.), 10.5 hours (Sat.)

→ large centrifugal force

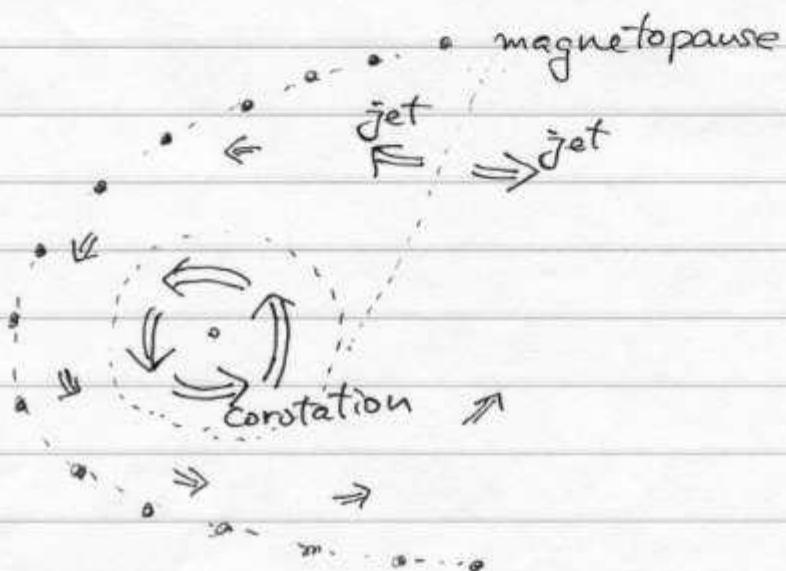
aurora ... magnetosphere-ionosphere coupling.
radio emission source

* Saturn dipole axis
aligned with rotation
axis

Force balance



View from rotation axis (Jupiter)



periodic jet \rightarrow substorm? reconnection?
on morning side \downarrow
internally driven substorm?

Plasma source from satellites

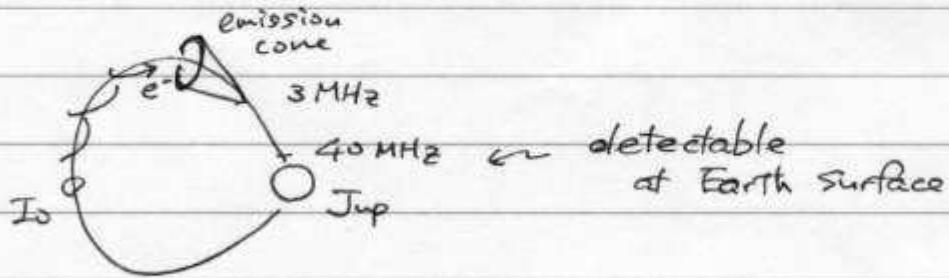
Io (Jup.) ... volcanism, sulfur-rich plasma torus

Enceladus (Sat.) ... water ice, water- ion products

Radio emission

- Auroral emission (cyclotron maser)
- radiation belt (synchrotron)

Jupiter aurora emission



Uranus and Neptune

Magnetized, gas planets

Surface mag. field 23 000 nT (U), 14 000 nT (N)

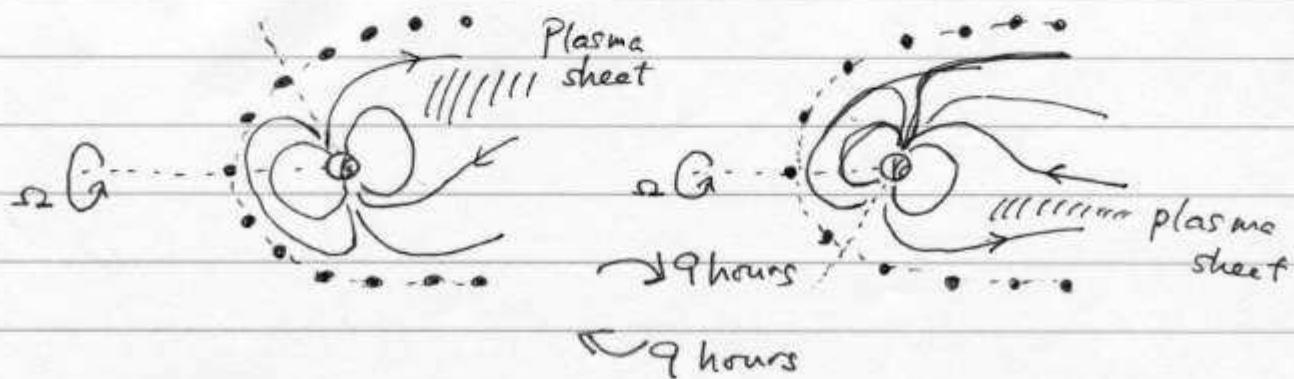
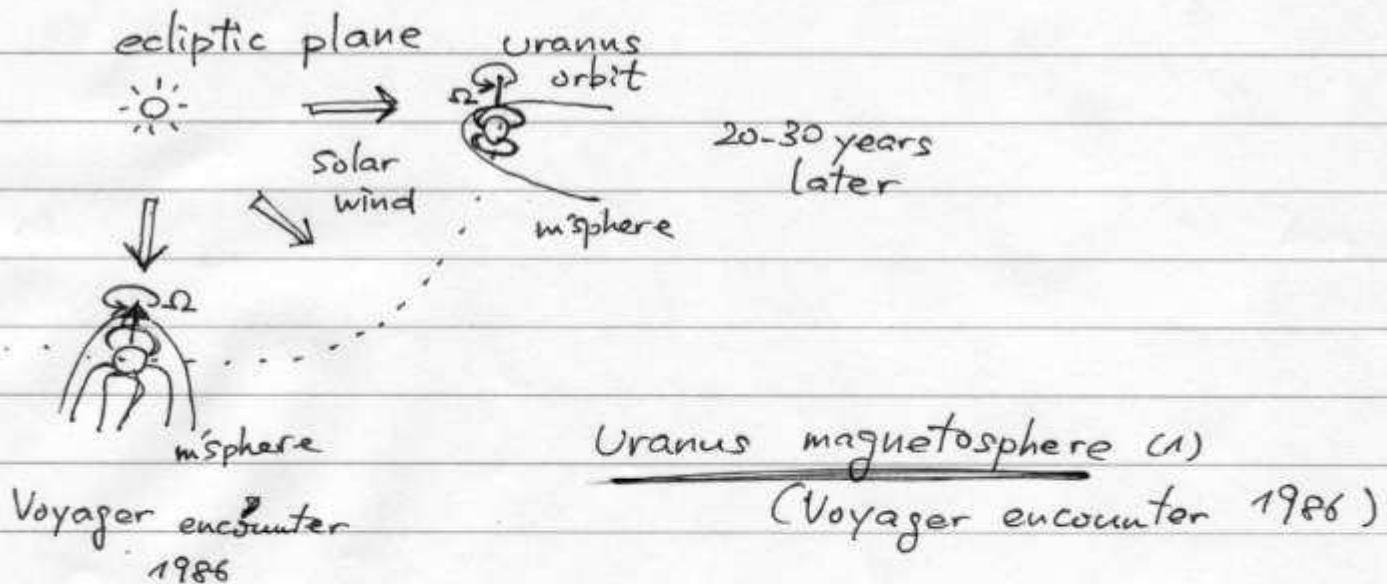
magnetopause distance $25 R_{\text{surf}}$ (U) $26 R_{\text{surf}}$ (N)

Large tilt angle (rotation axis - dipole axis)

59° (U), 47° (N)

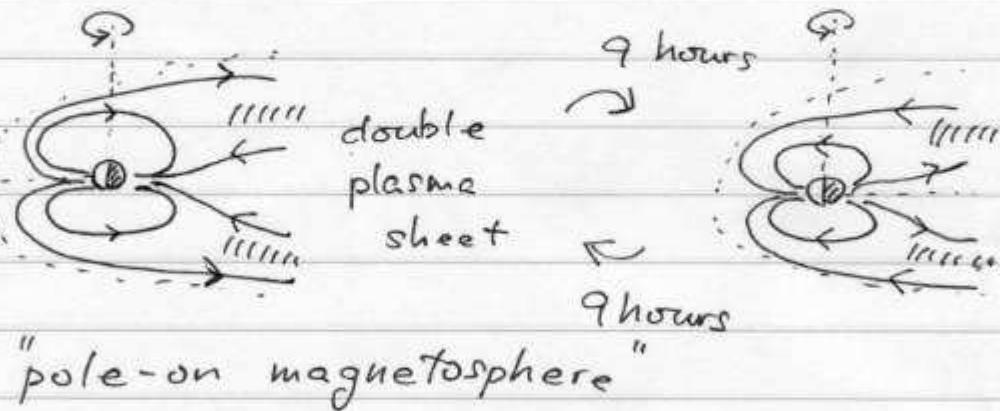
Uranus

Rotation axis almost in the orbital plane



Uranus magnetosphere (2)

20-30 years later



Neptune

